# A direct approach to detection and attribution of climate change

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Joint work with:

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- Reto Knutti (Institute for Atmospheric and Climate Science, ETH Zürich)
- Erich Fischer (Institute for Atmospheric and Climate Science, ETH Zürich)
- Nicolai Meinshausen (Seminar for Statistics, ETH Zürich)
- Guillaume Obozinski (Swiss Data Science Center)

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The climate system				



- The Earth is a complex system with highly nonlinear and unknown feedbacks between the components of the climate (atmosphere, ocean, land and ice)
- Internal (natural, unforced) climate variability: processes that occur naturally within the climate (El-Niño, North Atlantic Oscillation)
- Forced climate variability: external forcings (sun, volcanic activity, human influence)

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#### The climate system



- Regionally, the internal (natural) variability is large enough to mask climate changes for possibly many decades to come.
- **Open questions**: What changes will happen on regional scales and on shorter timescales? How strongly will the Earth's temperature respond to increasing CO2 levels?

#### Detection and attribution (D&A) of climate change



D&A studies the causal links between external drivers of climate change and observed changes in climate variables.

- **Detection** aims to find if there is a change in the observations that cannot be explained by internal (natural) climate variability alone.
- Attribution tries to assign the detected change to a particular external forcing or a combination of forcings.

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#### Traditional D&A

Global mean temperature anomaly w.r.t. 1850-1900



$$X_{obs} = f(X_{ant}, X_{nat}, \varepsilon_X)$$

- Unsupervised fingerprint (EOF) extraction
- Empirical Orthogonal Functions (EOF) ⇔ Principal Component Analysis (PCA)

$$\hat{X}_{obs}^{EOF} = \alpha_{ant} X_{ant}^{EOF} + \alpha_{nat} X_{nat}^{EOF}$$

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Direct D&A				

x = climate variable (temperature, humidity, precipitation) y = external forcing (anthropogenic, solar, volcanic)

$$\hat{eta} = \operatorname*{argmin}_{eta} \mathbb{E}_{(x,y)\sim P} \left[ l(y, f_{eta}(x)) 
ight]$$
 $\hat{y}_{new} = f_{eta}(x_{new})$ 

• Supervised fingerprint  $(\beta)$  extraction

• The predicted forcing  $\hat{y}$  used as a test statistic for D&A





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#### Detection and attribution as hypothesis testing

#### Detection:

- Null hypothesis: absence of an externally forced climate change
- The predicted forcing  $\hat{y}$  is indistinguishable from internal variability, e.g., not significantly different from zero

#### Attribution:

• If the true forcing y lies within the confidence intervals of the predicted forcing  $\hat{y}$ , then we attribute the detected change to the respective forcing

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#### Local vs. global daily weather



Signal vs. noise problem

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#### Detection of externally forced climate change



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#### Daily detection in global weather

• Assess whether externally forced climate change can be detected on shorter timescales (daily) if the detection is based on a *global* spatial pattern

nature	LETTERS
climate change	https://doi.org/10.1038/s41558-019-0666-7

## Climate change now detectable from any single day of weather at global scale

Sebastian Sippel<sup>12,3\*</sup>, Nicolai Meinshausen<sup>2</sup>, Erich M. Fischer<sup>1</sup>, Enikő Székely<sup>14</sup> and Reto Knutti<sup>1</sup>

For generations, climate scientists have educated the public that 'weather is not climate', and climate change has been framed as the change in the distribution of weather that slowly emerges from large variability is now in unchated terriweather when considered globally is now in unchated territories and the state of the state of the state of the observed temperature and molisticmer, we detect the fingerprint of externally driven climate change, and conclude that Earth as a whole is warming. Our detection approach involves statisti-

We start with a simple example to illustrate the difference in warming experienced locally and globally (Fig. 1). The past decade (2009–2018) has been on average 0.7°C warmer than an earlier period (1931–1980). Locally, decaseonalized daily temperature anomalies lactuate due to internal weather-related variability with season. This substantial variation implies that despite an overall warmer climate, cold anomalies or even cold records can still occur and are to be expected?. However, at the global scale, weather-

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#### Daily detection in global weather



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#### Daily detection in global weather

- The fingerprint of climate change is detected from any single day in the observed global record since early 2012, and since 1999 on the basis of a year of data
- While changes in weather locally are emerging over decades, global climate change is now detected instantaneously



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#### Distributional robustness



 $F_1 = \text{solar forcing}$   $F_2 = \text{volcanic forcing}$  $F_3 = \text{anthropogenic forcing}$ 

$$y = F_3$$

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#### Distributional robustness



- $F_1 =$ solar forcing  $F_2 =$ volcanic forcing
- $F_3 = anthropogenic forcing$

 $y = F_3$ 

 $F_1, F_2, F_3$  have a causal effect on x

$$x = g(F_1, F_2, F_3, \varepsilon_X)$$
$$y = f_\beta(x)$$
$$(x, y) \sim P$$

- Changes in  $p(F_1) \Rightarrow$  changes in p(x, y) even if p(y) remains the same
- Causal inference: interventions or perturbations

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#### Distributional robustness



**Goal**: Good prediction accuracy even under distributional changes of the external forcings (shift interventions).

Observational distribution  $(x, y) \sim P$ :

$$\hat{\beta} = \underset{\beta}{\operatorname{argmin}} \mathbb{E}_{(x,y)\sim P} \left[ I(y, f_{\beta}(x)) \right]$$

- $F_1 =$ solar forcing
- $F_2 =$  volcanic forcing
- $F_3 = anthropogenic forcing$

Class of distributions  $(x, y) \sim Q$  where  $Q \in Q$ :

$$\hat{eta} = \operatorname*{argmin}_{eta} \sup_{oldsymbol{Q} \in \mathcal{Q}} \mathbb{E}_{(x,y) \sim oldsymbol{Q}} \left[ l(y, f_{eta}(x)) 
ight]$$

 $y = F_3$ 

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Anchor regression

Anchor regression estimator [Rothenhäusler et al. (2019)]:

$$\hat{\beta}^{\gamma} = \underset{\beta}{\operatorname{argmin}} \|(I_n - \Pi_A)(Y - X\beta)\|_2^2 + \gamma \|\Pi_A(Y - X\beta)\|_2^2 \qquad \qquad y = F_3$$
$$A = [F_1, F_2]$$

•  $\Pi_A \in \mathbb{R}^{n \times n}$  is the matrix that projects on the column space of  $A \in \mathbb{R}^{n \times q}$ , i.e.,  $\Pi_A = A(A^T A)^{-1} A^T$ 

- $\gamma =$  "causal" regularization parameter that gives the strength of the intervention on the anchor variable A
- The causal regularization encourages orthogonality (uncorrelatedness) of the residuals with the anchor variable.

 $F_1$ 

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#### Anchor regression

$$\hat{\beta}^{\gamma} = \underset{\beta}{\operatorname{argmin}} \|(I_n - \Pi_A)(Y - X\beta)\|_2^2 + \gamma \|\Pi_A(Y - X\beta)\|_2^2$$

$$\begin{split} \gamma &= 0 \quad \Rightarrow \quad \hat{\beta}^{0} = \operatorname{argmin}_{\beta} \| (I - \Pi_{A})(Y - X\beta) \|_{2}^{2} & \text{Partialling out A (PA)} \\ \gamma &= 1 \quad \Rightarrow \quad \hat{\beta}^{1} = \operatorname{argmin}_{\beta} \| Y - X\beta \|_{2}^{2} & \text{Ordinary Least Squares (OLS)} \\ \gamma &\to \infty \quad \Rightarrow \quad \hat{\beta}^{\infty} = \operatorname{argmin}_{\beta} \| \Pi_{A}(Y - X\beta) \|_{2}^{2} & \text{Instrumental variables (IV)} \Rightarrow \text{Causality} \end{split}$$

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Anchor regression				

Let  $X \in \mathbb{R}^{n \times p}$ ,  $Y \in \mathbb{R}^{n}$ ,  $\beta \in \mathbb{R}^{p}$ , n = number of samples, p = dimensionality

$$\begin{split} \tilde{X} &= (I_n - \Pi_A)X + \sqrt{\gamma}\Pi_A X \\ \tilde{Y} &= (I_n - \Pi_A)Y + \sqrt{\gamma}\Pi_A Y \end{split}$$

 $\hat{\beta}^{\gamma} = \mathop{\rm argmin}_{\beta} \|\tilde{Y} - \tilde{X}\beta\|_2^2$ 

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Anchor regression				

Let  $X \in \mathbb{R}^{n \times p}$ ,  $Y \in \mathbb{R}^n$ ,  $\beta \in \mathbb{R}^p$ , n = number of samples, p = dimensionality

 $\tilde{X} = (I_n - \Pi_A)X + \sqrt{\gamma}\Pi_A X$  $\tilde{Y} = (I_n - \Pi_A)Y + \sqrt{\gamma}\Pi_A Y$ 

$$\hat{eta}^\gamma = \mathop{\mathsf{argmin}}_eta \| ilde{Y} - ilde{X} eta \|_2^2$$

• Ridge regularization allows to handle the multicollinearity of the predictors

$$\hat{\beta}^{\gamma} = \operatorname*{argmin}_{\beta} \|\tilde{Y} - \tilde{X}\beta\|_2^2 + \lambda \|\beta\|_2^2$$

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Data				

#### Climate simulations:

- CMIP5 (Coupled Model Intercomparison Project)
- control runs (climate simulations with no external forcings, only internal variability)
- Representative Concentration Pathways (RCP) (climate simulations with all external forcings)



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#### Data

- Climate simulations
  - 42 RCP 8.5 and 40 control runs
  - Temperature and precipitation
- Temporal resolution
  - annual (1870-2100)
  - $n = 82 \times 231 = 18,942$  samples
- Spatial resolution
  - $p = 144 \times 72 = 10,368$  dimensions



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Radiative forcing				



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Cross validation				



Cross validation (anchor regression  $\gamma = 10000$ ): PR -- ANTHRO forcing [rcp85] (1870 - 2100)

Model-wise splitting for training - testing - folds.

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#### Temperature - y = ANT, A = VOL



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#### Precipitation - y = ANT, A = VOL



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#### Attribution of externally forced climate change

• If the true forcing y lies within the confidence intervals of the predicted forcing  $\hat{y}$ 



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Summary				

- Introduced a novel statistical learning approach for D&A that fits into the framework of supervised methods
- We can now detect the signal of externally forced climate change in any single day since 2012 when considering the weather globally
- Distributional robustness protects us against future distributional changes allowing us to do attribution



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Future work				

- Apply the direct D&A approach to observations, other climate variables, multivariate relationships (temperature and precipitation)
- Relationship between external forcings and the water cycle is more complex than temperature because of larger internal variability which can mask externally forced climate changes
- Daily detection has implications for extreme events attribution



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### Thank you!

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